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Author for correspondence:
Yongjian Ding,
E-mail: dyj@lzb.ac.cn

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A new global gridded glacier dataset based on the Randolph Glacier Inventory version 6.0

Yaojun Li[1,2,3], Fei Li[3,4], Donghui Shangguan[1] and Yongjian Ding[1,2,3]

[1]State Key Laboratory of Cryospheric Science, Northwest Institute of Eco-Environment and Resources, Chinese Academy of Sciences, Lanzhou, China; [2]Key Laboratory of Ecohydrology of Inland River Basin, Northwest Institute of Eco-Environment and Resources, Chinese Academy of Sciences, Lanzhou, China; [3]University of Chinese Academy of Sciences, Beijing, China and [4]Key Laboratory of Tibetan Environment Change and Land Surface Processes, Institute of Tibetan Plateau Research, Chinese Academy of Sciences, Beijing, China

Abstract

Gridded glacier datasets are essential for various glaciological and climatological research because they link glacier cover with the corresponding gridded meteorological variables. However, there are significant differences between the gridded data and the shapefile data in the total area calculations in the Randolph Glacier Inventory (RGI) 6.0 at global and regional scales. Here, we present a new global gridded glacier dataset based on the RGI 6.0 that eliminates the differences. The dataset is made by dividing the glacier polygons using cell boundaries and then recalculating the area of each polygon in the cell. Our dataset (1) exhibits a good agreement with the RGI area values for those regions in which gridded areas showed a generally good consistency with those in the shapefile data, and (2) reduces the errors existing in the current RGI gridded dataset. All data and code used in this study are freely available and we provide two examples to demonstrate the application of this new gridded dataset.

Introduction

Since the release of its first version in February 2012, the Randolph Glacier Inventory (RGI) has been a centerpiece of glaciological research and successfully applied in a range of glacier change studies, as well as in impact assessments of glacial change at global and regional scales (Pfeffer and others, 2014; Hock and others, 2019). The latest version (6.0) of the RGI was released in July 2017 (RGI Consortium, 2017). In it, glacier outlines and their accompanying attributes in RGI are provided as shapefiles (a vector format). In addition, a gridded glacier map is supplied at 0.5° spatial resolution in which zonal records of glacierized areas (in km^2) are stored in a plain-text .DAT file (RGI Consortium, 2017). However, there are significant differences in total glacierized area between the gridded and the shapefile data in RGI 6.0, at both global and regional scales (Table 1).

Climate change studies frequently rely on gridded datasets, which are easy to use and comprise an effective tool for various glaciological and climatic applications, including, but not limited to, glacier change assessments (e.g. Brun and others, 2017), glaciological characteristics (e.g. Scherler and others, 2018), assessing glacier response to climate change (e.g. Sakai and Fujita, 2017), hydrology-related research (Huss and Hock, 2018) and modeling glacier changes (e.g. Raper and Braithwaite, 2006; Shannon and others, 2019). Increasingly, global and regional gridded datasets at different resolutions, such as the outputs of Coupled Model Intercomparison Project phase 6 (CMIP6, Eyring and others, 2016), are emerging and gridded data are critical for linking between glaciers and the corresponding meteorological variables. In this study, we identify errors in the currently available global gridded glacier dataset and present a new global gridded glacier dataset based on the RGI 6.0 which used an alternative gridding method to eliminate the errors.

Data and methods

We obtained the RGI 6.0 shapefiles as input (original) data, as well as the 0.5° × 0.5° grid data for the (following) comparative analysis, both of which are freely available at GLIMS (https://www.glims.org/RGI/rgi60_dl.html). The steps to generate the gridded dataset are shown in Figure 1. First, we produced the global grid map at a given resolution which is referenced to the WGS84 datum. Second, we split the glaciers using the boundaries of the cell. Third, we re-projected all polygons in the cell to Mollweide projection and recalculated the area of each polygon in the cell, and then summed them up to get the total glacier area of the cell at last (Fig. 1). Compared to the previous method, which typically uses the center point of glaciers and attributes the area to the cells in which the center point is located (so-called CP method), we eliminated the error caused by the glaciers overlapping grids which is unavoidable in the CP method. Since the sources of glacier inventory outlines are remarkably diverse, we recommend using the method by Pfeffer and others (2014) to propagate uncertainty, as this eliminates the effort of the inventory source. The errors for each grid in the 0.5° dataset are provided in the Supplementary material.

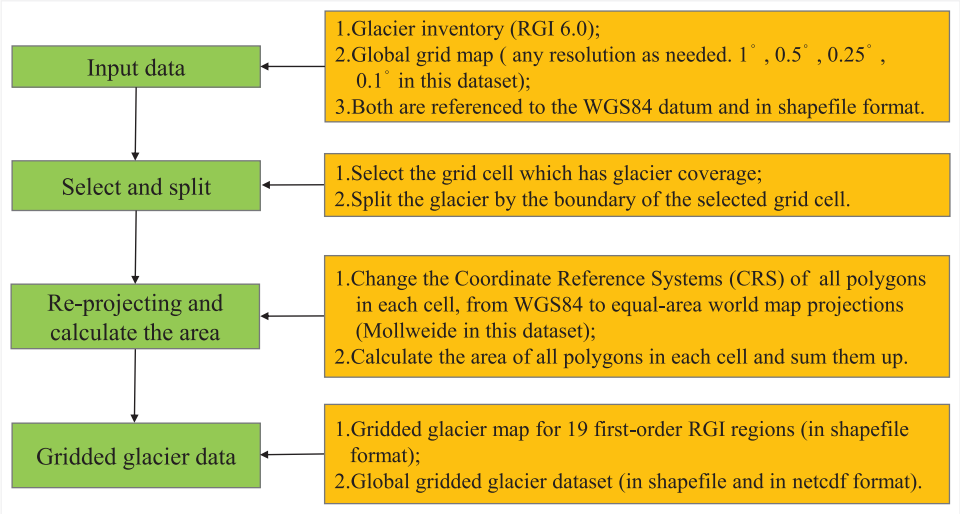


Fig. 1. Schematic representation of gridded glacier data.

Table 1. Area comparison between the shapefile and gridded map in RGI 6.0

ID	Region name	Shapefile area (km^2)	Gridded area (km^2)	Difference (%)
1	Alaska	86 725	86 691	−0.04
2	Western Canada and USA	14 524	14 558	0.23
3	Arctic Canada North	105 111	105 110	0.00
4	Arctic Canada South	40 888	40 888	0.00
5	Greenland Periphery	130 071/89 717[a]	89 717	0.00
6	Iceland	11 060	11 060	0.00
7	Svalbard	33 959	33 959	0.00
8	Scandinavia	2949	2949	0.00
9	Russian Arctic	51 592	51 592	0.00
10	North Asia	2410	2410	0.00
11	Central Europe	2092	2092	0.00
12	Caucasus and Middle East	1307	1306	−0.08
13	Central Asia	49 303	36 466	−26.04
14	South Asia West	33 568	15 121	−54.95
15	South Asia East	14 734	11 910	−19.17
16	Low Latitudes	2341	2341	0.00
17	Southern Andes	29 429	18 429	−37.38
18	New Zealand	1162	102	−91.22
19	Antarctic and Sub-Antarctic	132 867	93 507	−29.62
	Total	746 092/705 738	620 208	−12.12

[a]The total area of the region is 89 717 km^2 if glaciers strongly connected to Greenland ice sheet (~40 354 km^2) are excluded.

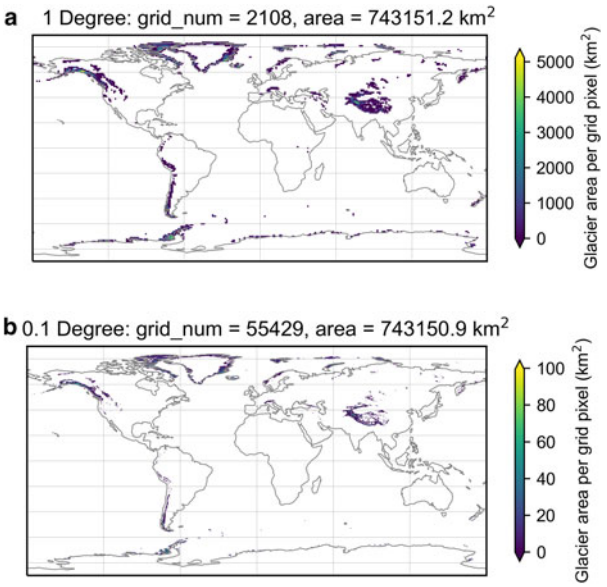


Fig. 2. Global gridded glacier map with two spatial resolutions. Grid_num means the number of glacierized grids.

Results and application

Here, we present the new global gridded glacier dataset based on the RGI 6.0. We provide four spatial resolutions of 1° × 1°, 0.5° × 0.5°, 0.25° × 0.25° and 0.1° × 0.1° as options to meet the requirements of different users (Fig. 2). It is noteworthy that glaciers strongly connected to the Greenland ice sheet (~40 354 km^2) are included into our gridded dataset. Our approach and code (data availability) are directly applicable to future regional and global glacier inventories. In addition, some attributes of glaciers could also be re-gridded by using our method. For example, we made a gridded dataset of supraglacial debris based on Herreid and Pellicciotti (2020) with a spatial resolution of 0.5° × 0.5° with this method (Fig. S1).

In Figure 3, we show a grid-based comparison between our dataset and the RGI gridded glacier dataset at 0.5° spatial resolution. As for the Greenland Periphery, given that glaciers strongly connected to the Greenland ice sheet (~40 354 km^2) are excluded from RGI gridded dataset, another version is created with our method (Fig. 1) for verification of the reliability of the new gridded glacier dataset (Fig. 3b). Our results are consistent with the RGI 6.0 gridded data for many regions, but we compute very

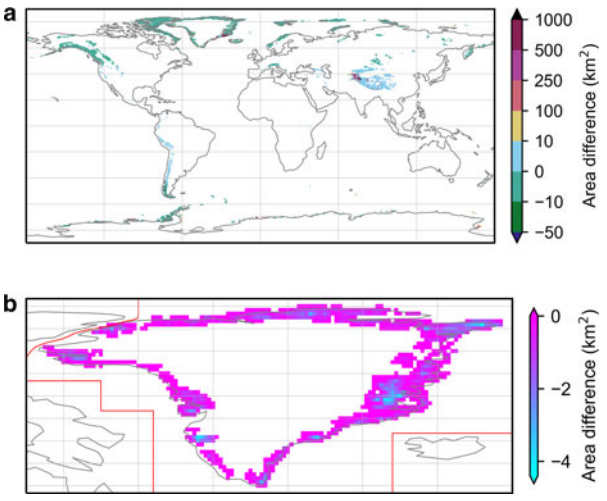


Fig. 3. (a) Grid-based comparison between the new dataset and RGI gridded glacier dataset at 0.5° spatial resolution globally. (b) Comparison of glaciers in Greenland periphery in two datasets when glaciers strongly connected to Greenland ice sheet are excluded.

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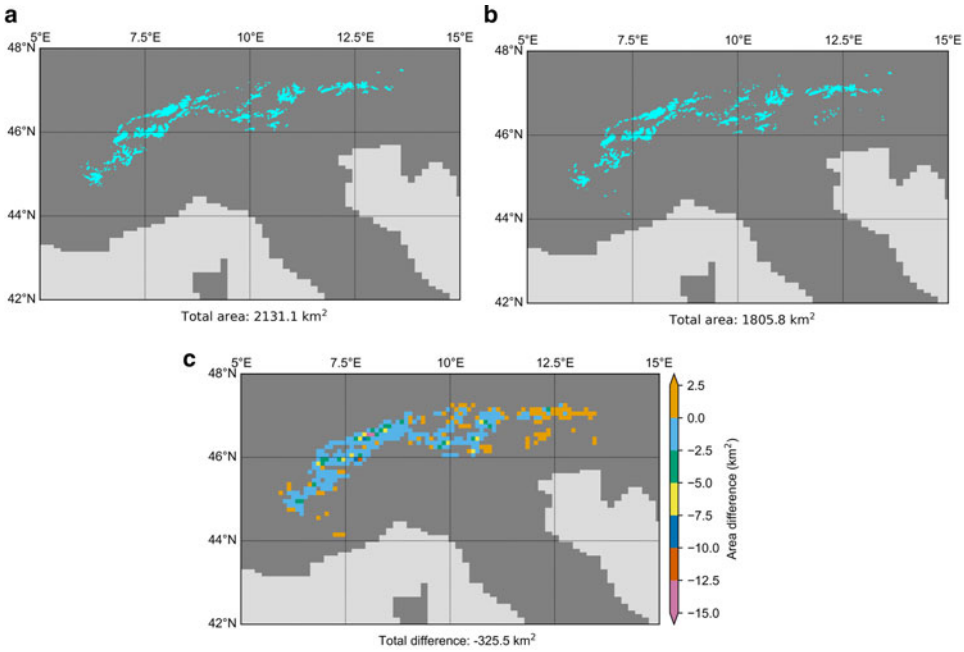


Fig. 4. (a) Glacier inventory of the Alps in 2003; (b) glacier inventory of the Alps in 2015 and (c) gridded area change (0.1° × 0.1°) between the two glacier inventories (in km^2).

Table 2. Area difference between the RGI 6.0 shapefile values and the new gridded dataset values for each region

ID	Region name	1° area (%)	0.5° area (%)	0.25° area (%)	0.1° area (%)
1	Alaska	−0.38	−0.38	−0.37	−0.35
2	Western Canada and USA	0.07	0.07	0.01	−0.06
3	Arctic Canada North	−0.62	−0.62	−0.62	−0.62
4	Arctic Canada South	−0.51	−0.51	−0.51	−0.51
5	Greenland Periphery[a]	−0.54	−0.54	−0.54	−0.54
6	Iceland	−0.42	−0.42	−0.42	−0.42
7	Svalbard	−0.62	−0.62	−0.62	−0.62
8	Scandinavia	−0.41	−0.41	−0.41	−0.41
9	Russian Arctic	−0.61	−0.61	−0.61	−0.61
10	North Asia	−0.22	−0.22	−0.22	−0.22
11	Central Europe	−0.02	−0.02	−0.02	−0.02
12	Caucasus and Middle East	−0.01	−0.01	−0.01	−0.01
13	Central Asia	0.77	0.77	−0.56	0.28
14	South Asia West	1.16	1.16	0.21	−0.17
15	South Asia East	−3.76	−3.76	2.84	1.00
16	Low Latitudes	0.61	0.61	0.61	0.61
17	Southern Andes	−0.07	−0.07	−0.07	−0.07
18	New Zealand	0.01	0.01	0.01	0.01
19	Antarctic and Sub-Antarctic	−0.51	−0.51	−0.51	−0.51
	Total	−0.39	−0.39	−0.39	−0.39

[a]Greenland glaciers strongly connected to the ice sheet are included.

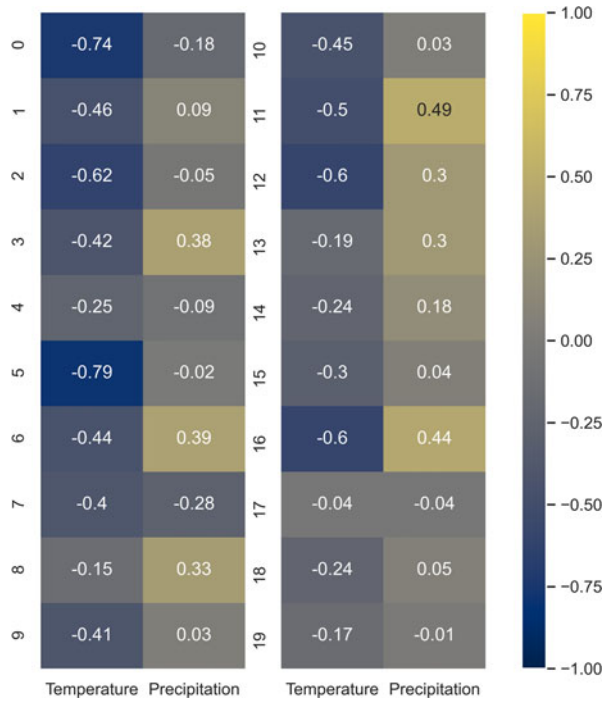


Fig. 5. Pearson correlation coefficients between annual mass balance and glacier area-averaged temperature and precipitation over the 1981–2016 period. Integers from 1 to 19 on the *y*-axis corresponding to the ID of the 19 RGI regions as defined in Table 1. The number 0 represents correlation coefficients on the global scale.

different gridded areas for High Mountain Asia, Southern Andes, New Zealand and Antarctic and Sub-Antarctic regions, where the RGI shapefile and gridded areas differ greatly to one another. Our dataset greatly reduced errors existing in the RGI gridded glacier map (Table 2). The accuracy of our gridded dataset depends critically on the selection of map projection and the method for area calculation. Since there is no previous study which utilized glacier polygons split by the boundaries of the cell for reference, we compared the area of each glacier provided by the RGI Consortium and the recalculated area to assess the map projection and the calculation method we used is reasonable or not (Fig. S2). We found that the recalculated area of each glacier was ~0.3% smaller than the RGI-provided area. However, 0.3% is considerably smaller than the uncertainty of the glacier extent in RGI (~5%, Pfeffer and others, 2014).

We demonstrate the application of our gridded dataset with two examples. First, we use two glacier inventories for the alps (Paul and others, 2011, 2020) to demonstrate the utility for glacier area change assessments. It is always hard to perform a glacier-by-glacier area change assessment between the latest and the earlier inventories due to the inhomogeneous interpretation of glacier extents, but a cell-by-cell comparison could reveal the spatial variability of glacier area changes (Fig. 4). The uncertainty in area change for each grid could be assessed following the law of error propagation:

$$\frac{\sqrt{e_i^2 + e_f^2}}{2}$$

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where e_i means the grid error in the initial state and e_f means the grid error in the final state, both of which could be calculated using the method by Pfeffer and others (2014). Second, we use this gridded dataset to determine correlation (Pearson coefficients) between glacier variations and climate changes on global scale for the period 1981–2016 (Fig. 5). The annual mass balance is derived from the latest compilation dataset of direct and geodetic observations developed by Zemp and others (2019). By using the method of area weighted average (Li and others, 2019) which is dependent on the glacierized areas (in km^2) in each cell, the annual climate variable series (including temperature and precipitation) for each first-order region can be derived from the ERA5 dataset with $0.5° \times 0.5°$ resolution (Hersbach and others, 2020). The results show that there is a higher level of correlation between mass balance and temperature than between mass balance and precipitation (Fig. 5).

We acknowledge that addressing the issue of gridded glacier maps in the RGI 6.0 is not technically difficult. However, it is an academic issue that must be overcome in order to provide more accurate datasets for glaciological research.

Conclusions

There are significant differences in the total glacier area between the gridded data and the shapefile data in the RGI 6.0 at both global and regional scales. Based on the RGI 6.0, we developed a new global gridded glacier dataset which reduced the errors of the RGI gridded glacier map. Moreover, we demonstrated the application of this gridded dataset with two examples.

Supplementary material. The supplementary material for this article can be found at https://doi.org/10.1017/jog.2021.28.

Data. The datasets and the code can be downloaded from https://github.com/rylanlee/RGI-Gridded.git and is linked on https://www.glims.org/RGI/rgi60_dl.html.

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